On the possible connection between cosmic rays and clouds Anatoly Erlykin^{*}, Gyula Gyalai[†], Karel Kudela[†], Terry Sloan[‡] and Arnold Wolfendale[§]

*P.N.Lebedev Physical Institute, Moscow, Russia [†]Institute of Experimental Physics, Kosice, Slovakia [‡]Department of Physics, Lancaster University, Lancaster, UK [§]Department of Physica, Durham University, Durham, UK

Abstract. Various aspects of the connection between cloud cover (CC) and cosmic rays (CR) are analysed. We argue that the anticorrelation between the temporal behaviour of low (LCC) and middle (MCC) clouds evidences against the causal connection between them and CR. Nevertheless, if a part of low clouds (LCC) is connected and varies with CR, then its most likely value averaged over the Globe should not exceed 20% at the two standard deviation level.

Keywords: correlation, cosmic rays, clouds

I. INTRODUCTION

A correlation between CR intensity and global LCC was observed for the first time more than 10 years ago [1], [2] and led to a new direction in science - cosmoclimatology [3]. It is based on the concept of a causal relationship between CR and CC. Some arguments against this causality were presented in [4] . The purpose of this study is to continue further an analysis of possible reasons for the observed correlation between CR and LCC.

II. INPUT DATA

As the input CC data, we took the same observations with weather satellites (ISCCP project) that were used in [1], [2]. We analysed sky fractions covered by clouds, averaged over observation months (D2 series). In compliance with the ISCCP cloud height classification made according to the pressure at their upper boundary, clouds were separated into low (LCC,>680 hPa), medium (MCC, 440-680 hPa) and high (HCC, <440hPa). Because of the ongoing discussion on the calibration quality of ISCCP radiometers after 1996 [5] we used both the data obtained earlier, during the 22nd cycle of solar activity (July 1986 - December 1995), and the complete set of data, till 2005. For comparison of CC and CR variations we used the data from several neutron monitors of the worldwide network (Thule, Apatity, Moscow, Climax, Huancayo).

In an analysis of the latitude dependence of CR and LCC variations we split the entire latitude range (from -90° to 90°) into nine equal intervals of 20° width. We also analysed the dynamics of global (i.e. globe-averaged) CC. To better reveal the CC variations of non-trivial origin, we subtracted the winter-summer seasonal variations from the temporal curves, although in some cases seasonal variations were considered as well.

III. RESULTS

A. Long-term variations and the fraction of LCC correlated with CR intensity



Fig. 1. CR and LCC variations and their correlation during the 22nd cycle of solar activity. (a) CR variation (counting rate in the Climax neutron monitor); (b) global LCC variation; (c) correlation of CR and LCC variations with respect to their average levels over the period 1986-1996 (long-term correlation), $b = 0.157 \pm 0.023$, r = 0.538, the best fit is $0.978 \pm 0.018x^{8.65}$; (d) correlation of CR and LCC variations with respect to the average temporal evolution curve - thin smooth lines in Figures 1a and 1b (short-term correlation), $b = -0.055\pm0.057$, r = -0.0904. The short-dash lines in all panels are the average levels over the period 1986-1996, the long-dash lines in Figures 1c and 1d are linear regression lines with a slope *b*, the thin smooth line in Fig.1c is the best power-law fit to the scatter plot, and *r* is the correlation coefficient.

Figure 1 shows the temporal evolution of (a) CR

intensity, (b) global LCC and (c) the correlation of CR and LCC deviations from their means. To illustrate the CR intensity evolution we took as a proxy the data of the Climax monitor, which is situated at a latitude of $39.4^{\circ}N$. The CR intensity fluctuations at other latitudes show qualitatively similar temporal behavior, although with different variation amplitudes. The CR and LCC deviations from their means are positively correlated. The linear regression slope is 0.157 ± 0.023 and the corresponding correlation coefficient is 0.538, which confirms the existence of positive correlation between CR and LCC that was found in [1], [2].

However, an attempt to understand what is the reason and what is the consequence here failed. We wanted to find a possible time shift between the CR and LCC variations by the method of least squares. However, it turned out that the sum of squared deviations has a flat broad minimum at a respective shift of the CR and LCC curves from -11 to +6 months. That is , one cannot say which variation is the cause and which is the consequence.

On the assumption that CR are indeed responsible for at least some part of the CC one can use the observed correlation to estimate this part. This estimation depends on the model for the relationship between CR and LCC. With a linear model $(\Delta/\langle\Delta\rangle = a + b(I/\langle I\rangle)^c$, where c = 1; Δ and I are the cloud coverage and CR intensity respectively; and $\langle \Delta \rangle$ and $\langle I \rangle$ are their mean values), the regression slope b = 0.157 gives the CC fraction related to CR, to be approximately 16%. Within two standard deviations this fraction should not exceed 20%. However, a least-squares estimation of the parameters a, b and c shows that the relation between CR and LCC is most likely non-linear. We obtained the values a = 0.978, b = 0.018 and c = 8.65 (thin solid line in Fig.1c). This shows that the most likely CC fraction related to CR does not exceed 2%.

The above conclusion is valid only if the models of relationship between CR and LCC are correct and the CR variations at the Climax latitude, $-40^{\circ}N$, adequately describe the global variation picture. For c < 1 the CC fraction which positively correlates with CR can be much higher. Unfortunately, because of the relatively small magnitude of CR and LCC variations one cannot reliably estimate the parameter c and thus evaluate more precisely the positively correlated CC fraction. Although the method of least squares is more adequate at c > 1, in the domain of available experimental data (Fig.1c) the behavior of the curves for different values of c is similar, as well as the corresponding sums of the squared deviations.

B. Short-term variations

Figure 1 and the analysis given in the previous section concern the total variations of CR and LCC relative to their average values. The main contribution to these variations is from the long-term variations related to the 11-year solar activity cycle. To reveal possible

correlations in short-time CR and LCC variations, we excluded the contribution from long-term variations. For this purpose we approximated the temporal evolution of CR and LCC by a fifth-order polynomial (smooth solid lines in Figures 1a and 1b) and calculated deviations from this approximation. Because we used the monthly averaged D2-series data, this analysis is related to monthly variations. We did not find any statistically significant correlations between the CR intensity and global LCC. The estimated regression line slope is $b = -0.055 \pm 0.057$ and the corresponding correlation coefficient is -0.090 (Fig.1d). This negative result is in indirect agreement with the absence of even shorter (few days) correlations (Forbush decreases, CR ground-level enhancements) pointed out in [4], [6].

C. Latitude dependence of correlations between CC and CR

Because the CR intensity depends on the latitude it is reasonable to analyse the variation of CC characteristics with latitude. Figure 2a shows the latitude dependence



Fig. 2. The latitude dependence of CC characteristics: (a) absolute values of LCC (open circles), MCC (full circles) and HCC (open stars). (b) LCC, MCC and HCC correlations with CR (Climax). Notations are the same as in (a).

of LCC, MCC and HCC. It is seen that there is a small minimum for LCC in the equatorial region, which could be connected with the reduction of CR intensity, but it is not confirmed by the local maxima in MCC and HCC. In polar regions, where the CR intensity is the highest, there is an opposite decrease of LCC, which apparently is connected with the dominant influence of the atmospheric conditions, eg. low temperatures. The highest LCC is in the southern latitude bands with the largest part of the area occupied by oceans, i.e. with a relatively large density of water vapor.

The altitude dependence of CC does not correspond to the altitude dependence of the CR intensity: a further bad feature. In most of the latitude bands, MCC and HCC are smaller than LCC, which is opposite to CR with their intensity rising with height. All this shows that even if there is a causal connection between CR and LCC, its character is more complicated than a direct and positive correlation.

We have already mentioned that the global LCC-CR correlation is positive: r = 0.538. Figure 2b shows the latitude dependence of the CC-CR correlation coefficient. Here again we used as a proxy of the CR temporal variations just the neutron counting rate at Climax. In spite of the latitude dependence of the CR variation amplitude, the value of the LCC-CR correlation coefficient does not depend on this amplitude due to the similarity of the temporal behavior of CR variations at different latitude bands. It is remarkable to note that in most latitude bands, MCC and HCC have negative correlation with CR in contrast to the positive LCC-CR correlation coefficient which was the main argument for the claimed causal CR-CC connection [1], [2].

D. Negative correlations of LCC and MCC

Figure 3 shows the latitude dependence of the sensitivity and correlation between MCC and LCC. The



Fig. 3. The latitude dependence of the sensitivity (slope b of the linear regression line) of MCC to LCC variations (full line) and their correlation coefficient r (dashed line).

sensitivity of one variable to another, according to the definition [7], is the derivative of the first variable on the second in log-coordinates. In our case the sensitivity is the slope of the linear regression line b in the MCC-LCC plot. One can notice two features:

(i) the sensitivity of MCC to LCC and MCC-LCC correlation coefficient are negative at all latitudes, which is another support of their global anticorrelation. The negative sensitivity of MCC to LCC is difficult to explain in the framework of the causal connection between CC and CR, since the rise of the CR intensity has to change CC similarly at all altitudes;

(ii) the highest negative sensitivity and the correlation between MCC and LCC is observed in tropical and subtropical regions $\ell = -30^{\circ}/+30^{\circ}$ as well as in southern latitude bands with the highest fraction of water: $\ell = -45^{\circ}/-65^{\circ}$.

IV. DISCUSSION

In our opinion the analysis performed here, as well as our previous arguments [4], gives grounds to assert that CR are not the dominant factor leading to CC formation. The negative correlation of LCC and MCC, most prominent in the tropics and subtropics and at the latitudes where one can expect excess water vapor formation, allows one to turn to the traditional picture describing the main reasons for cloud formation that are connected with solar activity.

Solar radiation increases together with the number of



Fig. 4. Temporal evolution of the surface temperature (a), HCC (b), MCC (c) and LCC (d) during two last cycles of solar activity. The dotted line in panel (a) is the average temperature in the 20th century. Dotted lines in panels (b)-(d) show the average CC level over the measurement period 1984-2005.

sunspots in the middle of the solar cycle. This radiation is strongest in the tropics and subtropics. Although the relative increase in radiation intensity is insignificant ($\sim 0.1\%$), it leads to an increase in the average groundlevel temperature and enhances vertical convective flows of heated air. Since the cloud height in the ISCCP experiment is classified according to the pressure at the upper cloud boundary, the convective lift of clouds to higher altitudes leads to a redistribution of the assigned altitudes. That is, it decreases LCC and increases MCC, which is reflected in the negative correlations of LCC and MCC. Thus, an increase in convective flows leads to a significant strengthening ($\sim 2\%$) of the effect of increased solar radiation.

Along with the periodic variations of the ground-level Earth's temperature, related to the periodicity of solar cycles, a systematic temperature increase due to global warming was observed in the last century. The warming was particularly strong during the last two decades. Figure 4a shows the average temperature growth, and Figures 4b-4d demonstrate the temporal evolution of HCC, MCC and LCC, respectively. One can see both periodic and systematic variations of MCC and LCC, which are negatively correlated. The data on the graphs suggest that the tendency of LCC reduction, which appeared after 1995 [5] is not due to malfunctioning of the weather satellite equipment but due to a global warming of the Earth's climate.

Summing up, one can state that with an increase in the ground-level Earth's temperature during solar activity maxima or general global warming, the average CC height becomes somewhat greater. This effect is most prominent in the tropics, subtropics and above the

surface of the ocean. The simultaneous reduction of LCC and of CR intensity is not evidence for a causal relationship between these two phenomena. They correlate due to the presence of a common driving force: changes in solar activity.

Acknowledgments

K.Kudela wishes to acknowledge the VEGA grant agency, project 2/7063/27. A.D.Erlykin acknowledges the support of the John C. Taylor Charitable Foundation and is grateful to E.A.Vashenyuk for granting access to the Apatity neutron monitor data and to S.P.Perov for useful discussions.

REFERENCES

- Svensmark H. and Friis-Christensen E., 1997, J. Atm. Solar-Terr. Phys., 59, 1225
- [2] Palle Bago, E. and Butler C.J., 2000, 41, 4
- [3] Svensmark H., 2007, News Rev. Astron. Geophys., 48, 118
- [4] Sloan T. and Wolfendale A.W.,2008, Environ. Res. Lett., 3, 024 001(6pp)
- [5] Marsh N. and Svensmark H., 2003, Space Sci. Rev., 107, 317
- [6] Kristjánsson et al., 2008, Atm. Chem. Phys., 8, 7373
- [7] Uchaikin V.V. and Ryzhov V.V., 1988, Stochastic theory of the transport of high-energy particles, Novosibirsk, Nauka (in Russian), p.144